



Groundwater Quality 2013

**Estimation of Contaminant Mass**  
**Discharges at Plume Control Planes**

Harald Klammler

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# Overview

- Concept of mass flux and discharge
- Problem of estimating discharges from point measurements → **UNCERTAINTY**
- Geostatistical approach for **POROUS AQUIFERS**
  - Spatial correlation (variogram)
  - Data skewness (non-Gaussian)
- Alternative approach for **FRACTURED ROCK**
  - Object oriented (linear fracture traces)

# What is Flux?

- Water (Darcy) flux:

$$q = Q/A \quad [L/T]$$

Q ... water discharge (volume per time)

A ... cross-sectional area

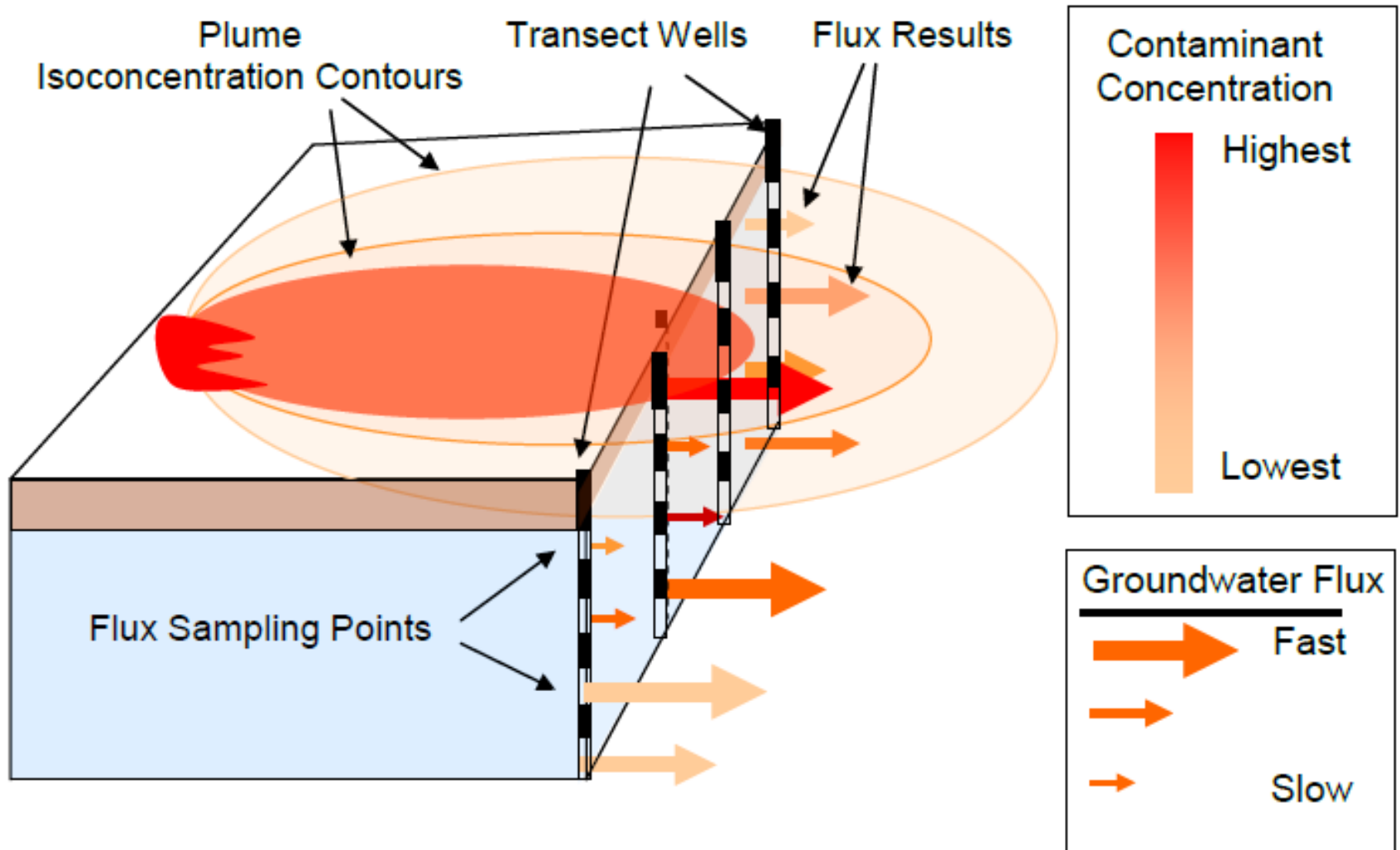
- Contaminant mass flux:

$$J = qC \quad [M/(L^2T)]$$

C ... contaminant concentration

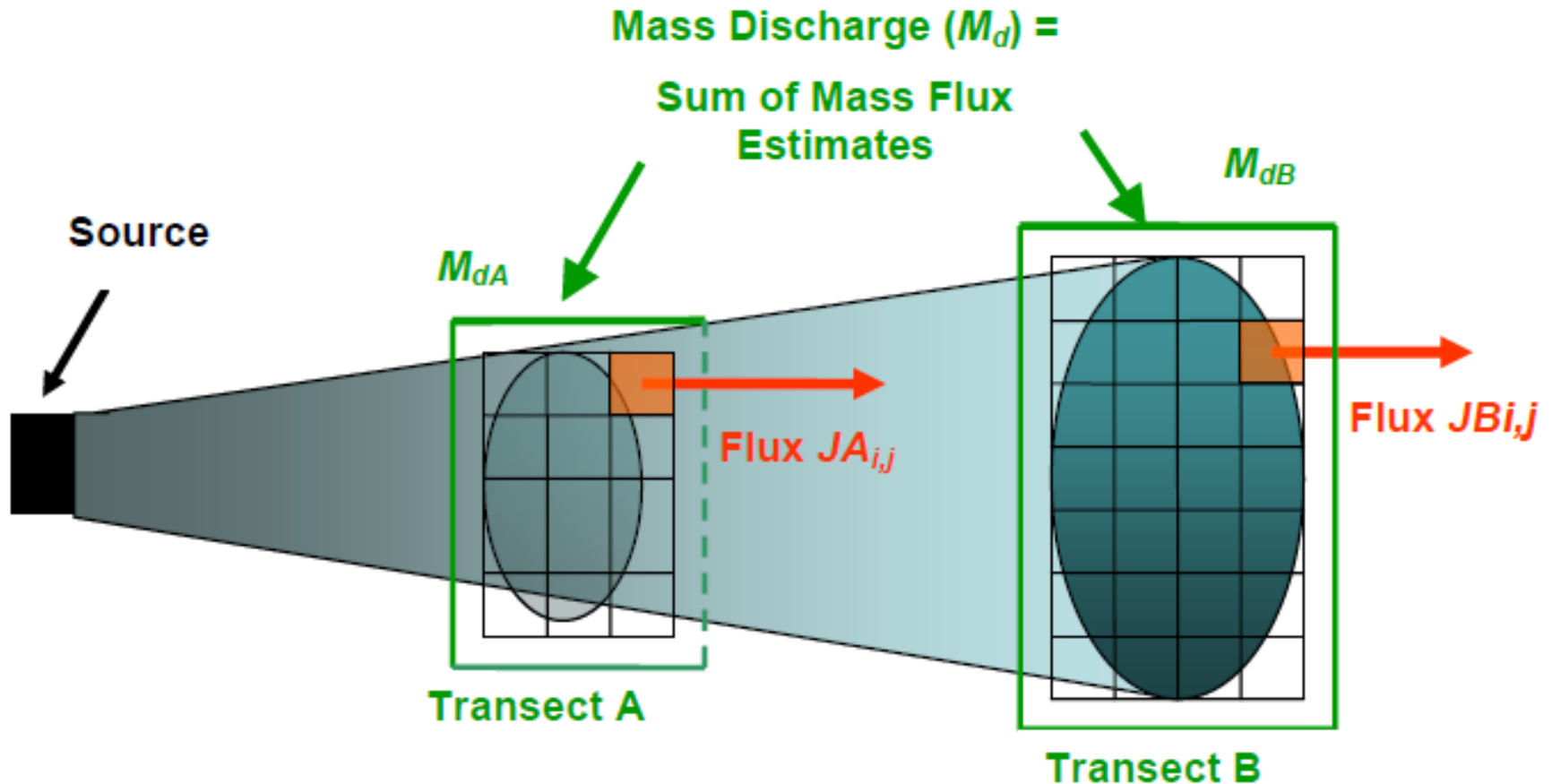
- **J is the mass of a contaminant passing a unit cross sectional area per unit time interval**

# What is Flux?



From ITRC (2010) "Use and Measurement of Mass Flux and Mass Discharge".

# What is Discharge?

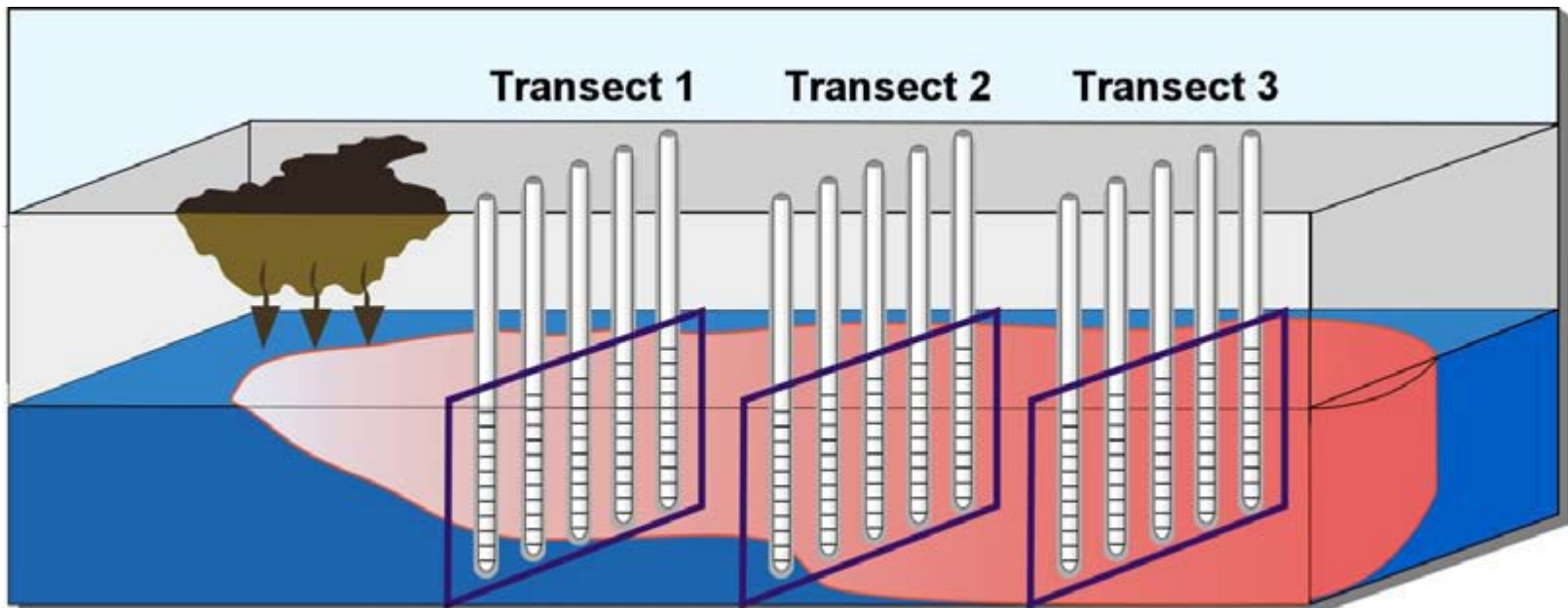


$JA_{i,j}$  = Individual mass flux measurement at Transect A

$M_{dA}$  = Mass discharge at Transect A (total of all  $JA_{i,j}$  estimates)

# Why Discharge?

- Quantify source strengths and plume attenuation rates
- Assess impact on potential receptors (e.g., supply well)
- Evaluate remediation performance or legal compliance



From ITRC (2010) "Use and Measurement of Mass Flux and Mass Discharge".

# Estimating Discharge

## (from point measurements)

- Complex methods incorporating different types of site information:
  - Heads, conductivities, fluxes, concentrations, etc.
- **FLUXES** are most “direct” information, because simple summation gives discharge
- Fluxes are only known at a **LIMITED NUMBER OF SAMPLING POINTS** in the control plane
- This requires **INTERPOLATION OF FLUXES** at unsampled locations before integration
- Interpolation introduces **UNCERTAINTY** due to random spatial variability of fluxes

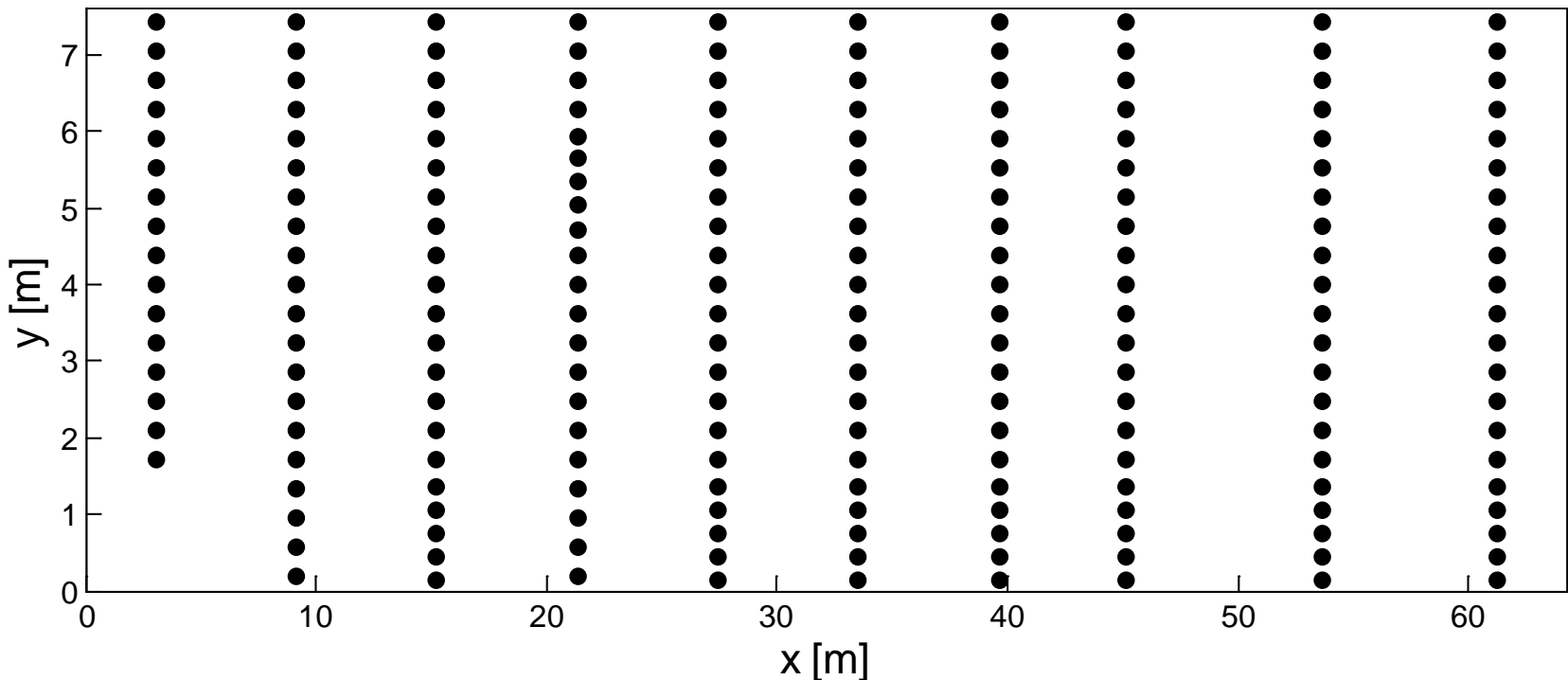
# Porous Aquifers

- Fluxes may be regarded as a continuous spatially random variable (**GEOSTATISTICS**)
- Spatial variability may be described by a **VARIOGRAM**
- **CONDITIONAL STOCHASTIC SIMULATION** can generate a large number of possible scenarios of flux distributions (discharges) across a control plane
- **PROBABILITY DISTRIBUTION OF DISCHARGE** to define confidence limits
- Simpler approximate methods have been developed based on an **EFFECTIVE NUMBER OF INDEPENDENT DATA**



# TCE plume, Ft. Lewis, WA

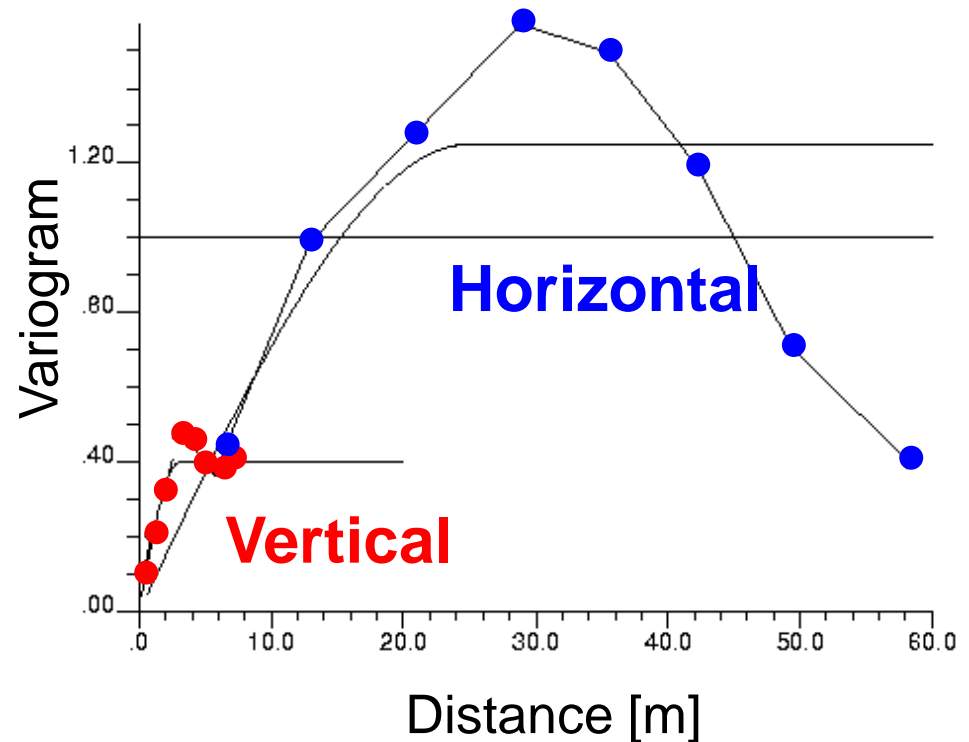
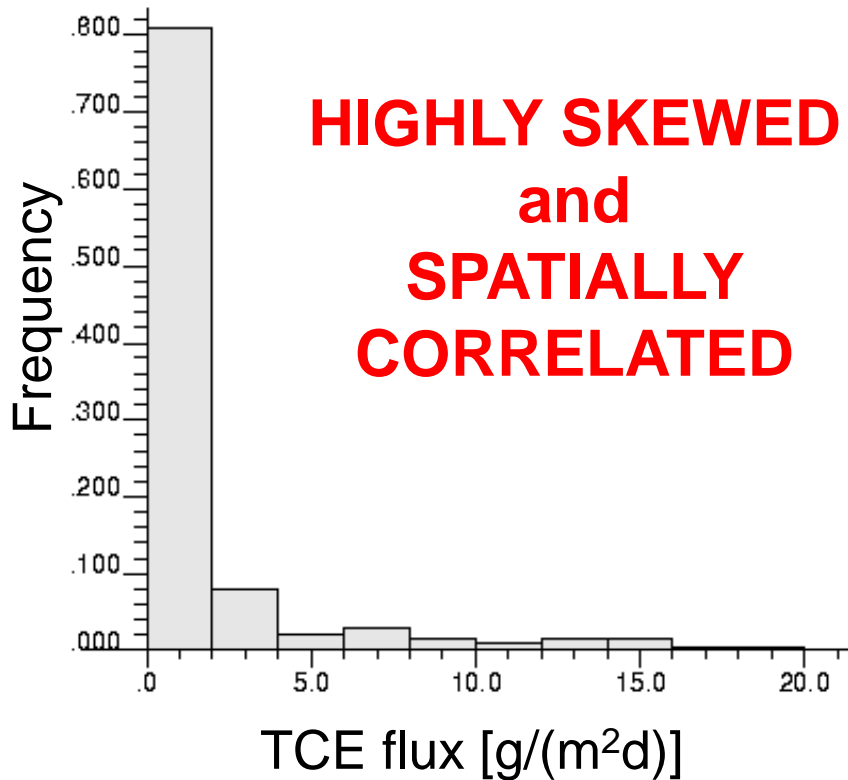
- Control plane with measured flux locations (possibly irregular)



(WRR, 2012, Contaminant discharge and uncertainty estimates from passive flux meter measurements)

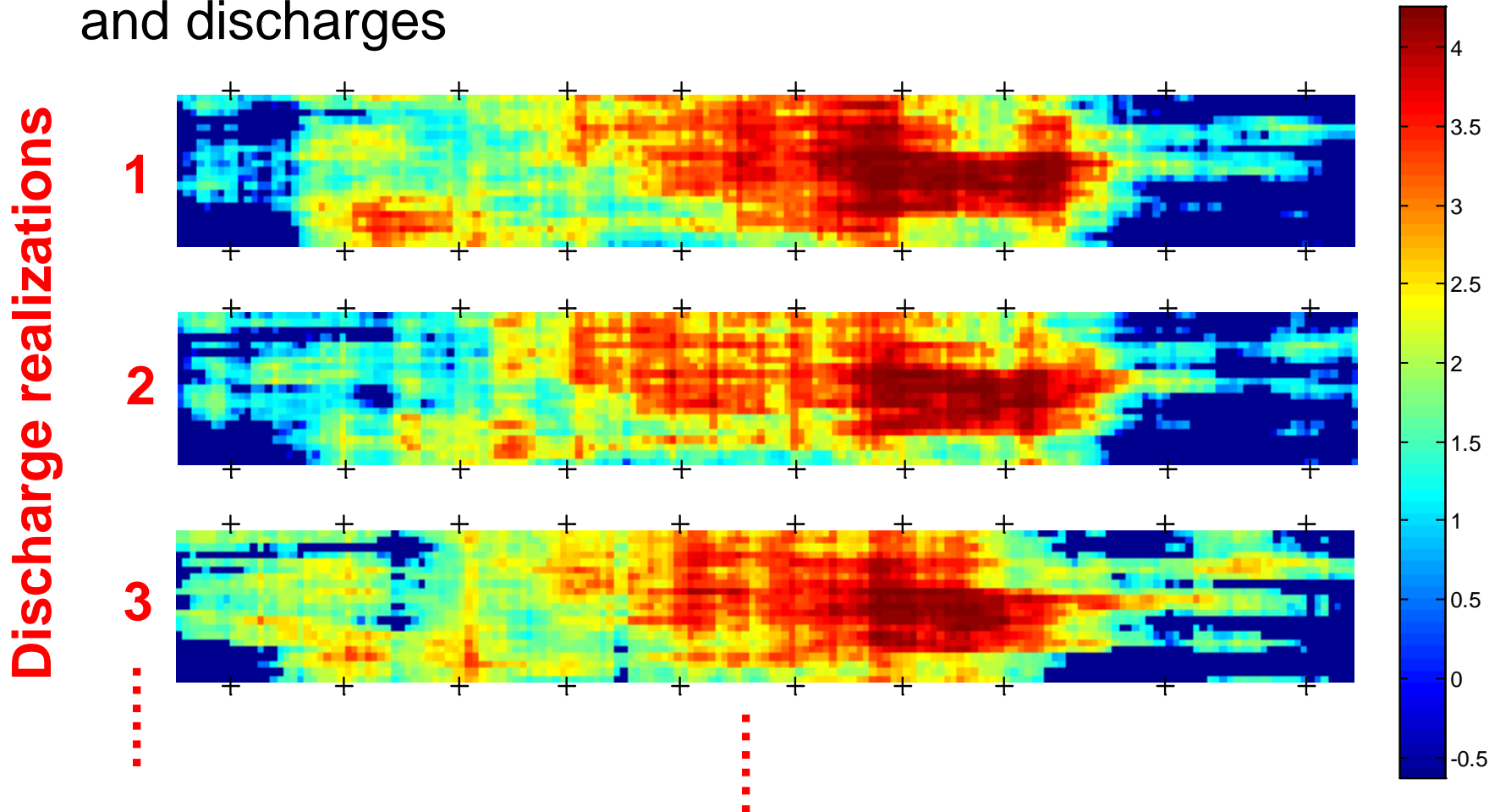
# TCE plume, Ft. Lewis, WA

- Flux histogram and variogram



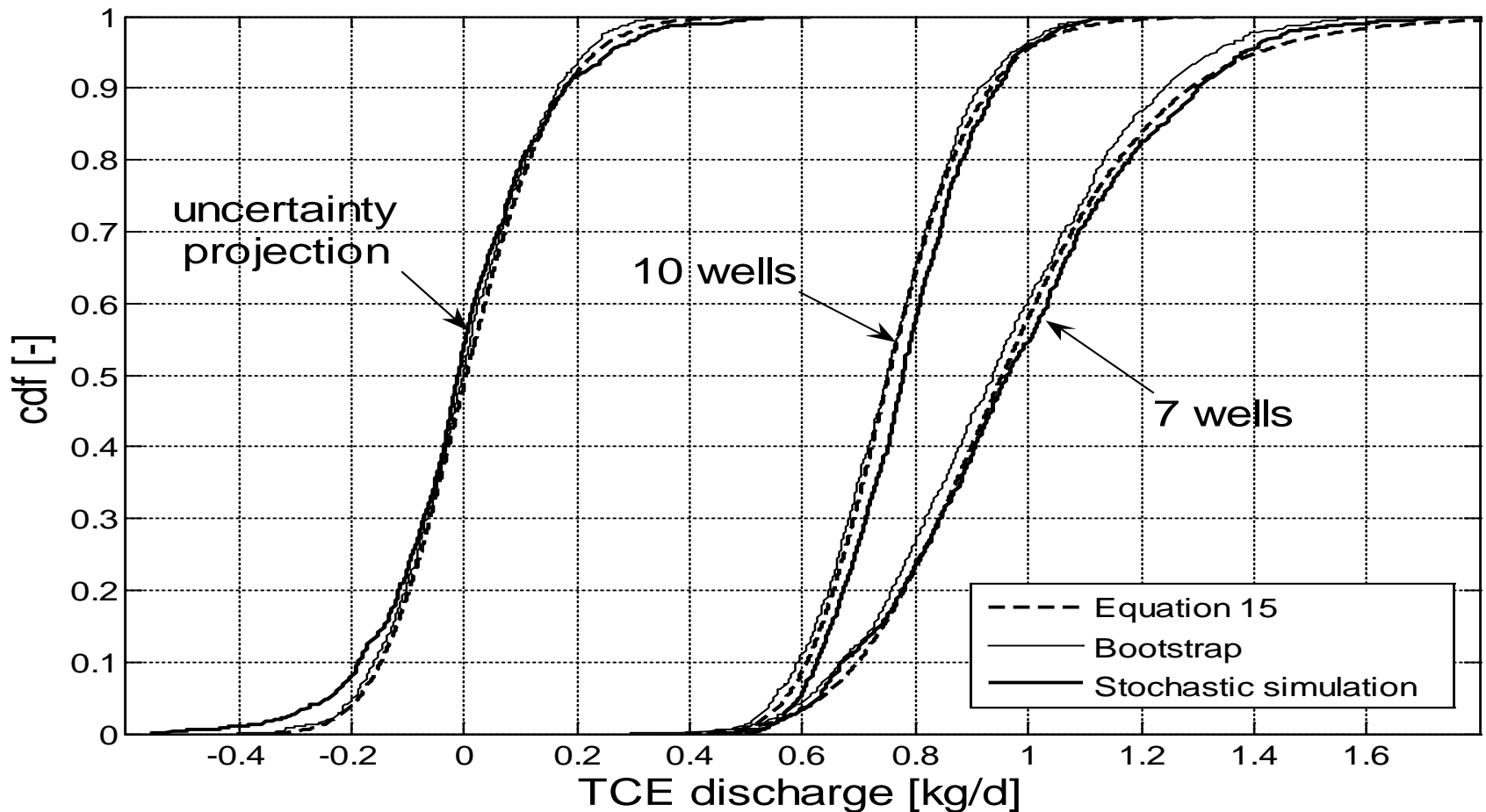
# TCE plume, Ft. Lewis, WA

- Three examples of simulated flux distributions (log-scale) and discharges



# TCE plume, Ft. Lewis, WA

- Cumulative distribution functions of TCE discharge

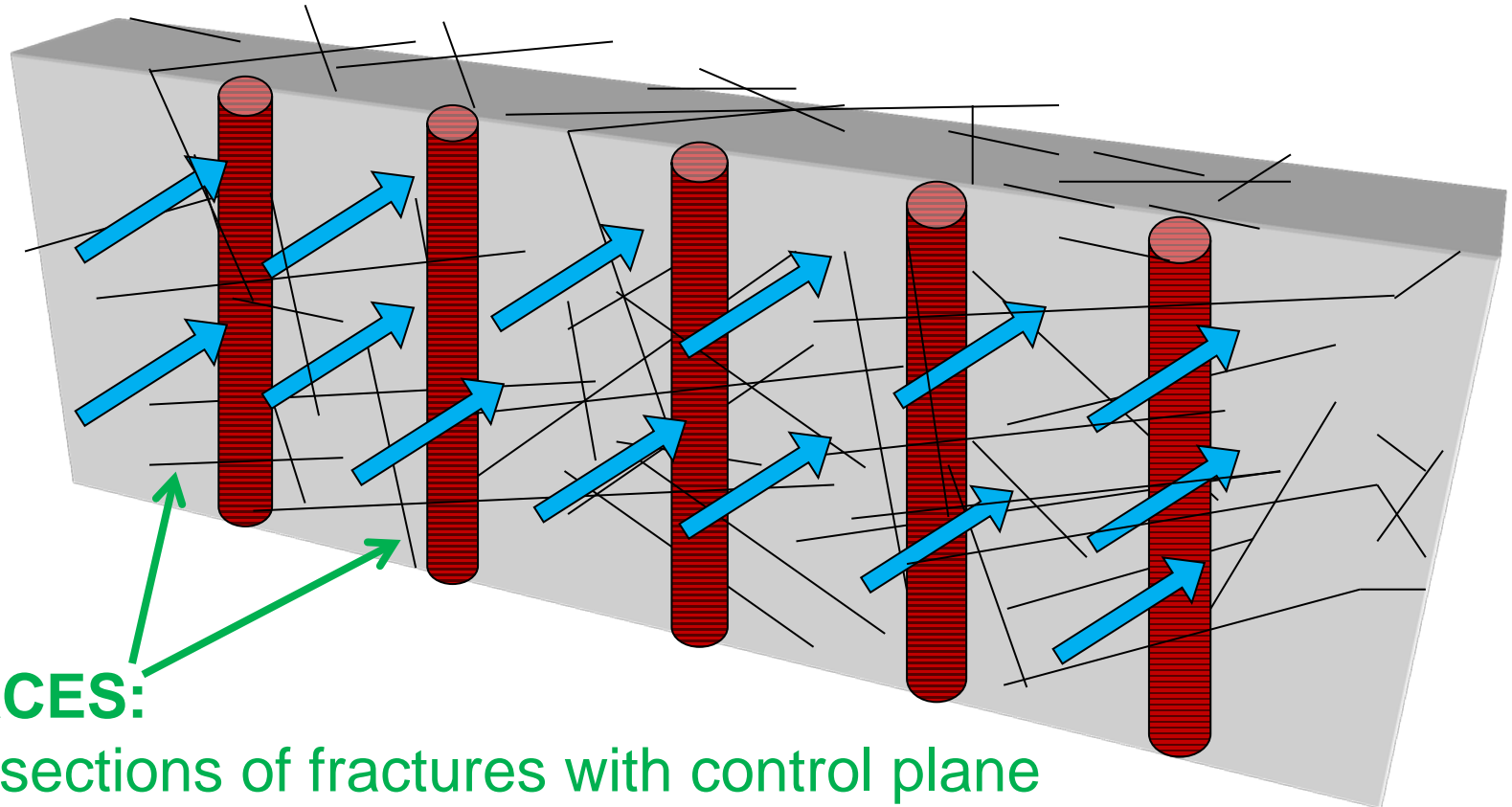


# Fractured Rock

- Flow occurs along fractures (predominantly)
- Fluxes are measured only where a borehole intersects a flowing fracture (assumed here to be plane)
- Sources of **UNCERTAINTY**:
  - Unknown fracture density
  - Random fracture locations, sizes and orientations
  - Random flux variability between fractures
  - Random flux variability within fractures
  - Random flux measurement errors

# Sampling Transect

- Control plane with observation wells intersecting flow and transport through fracture planes

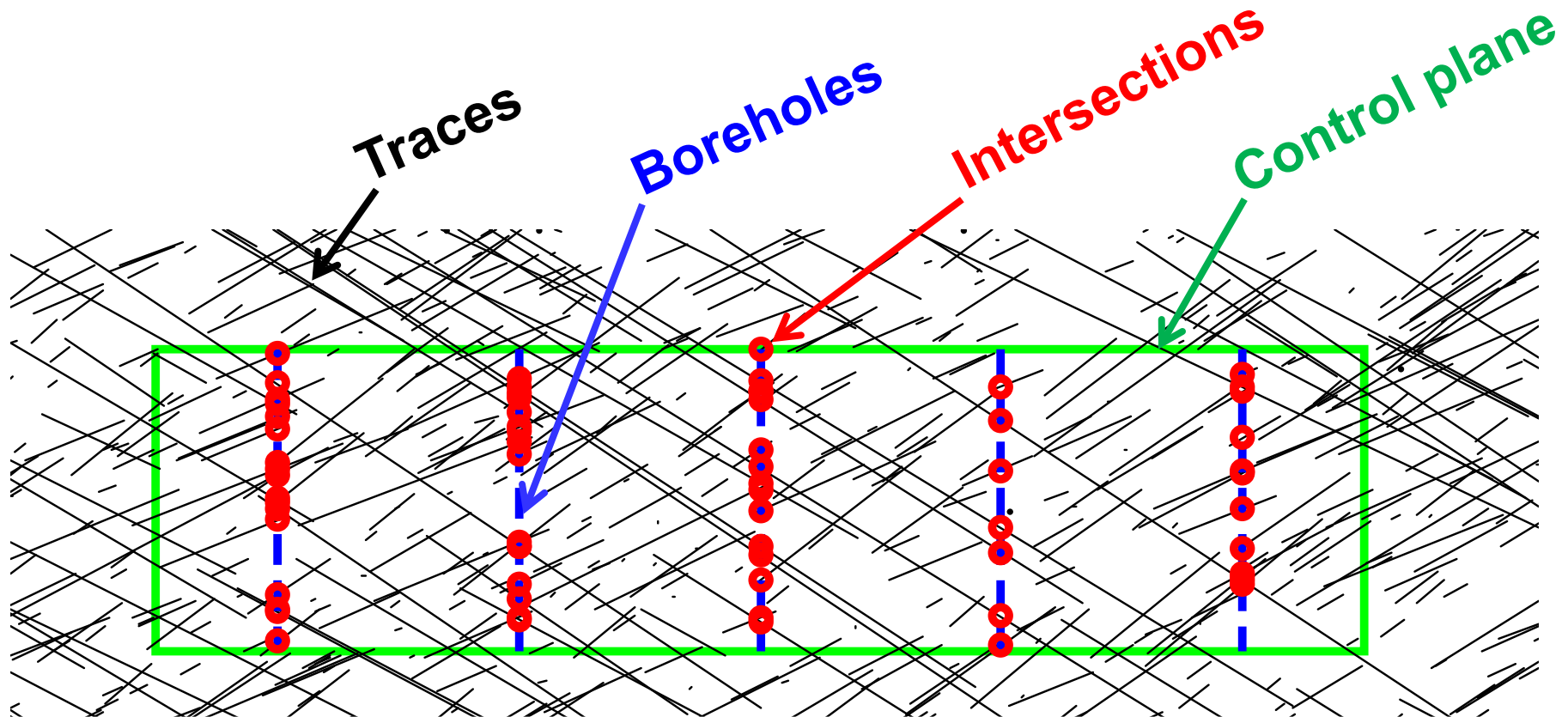


**TRACES:**

Intersections of fractures with control plane

# Conceptual Model

- Trace locations are assumed to follow Poisson process (complete randomness for a given density)



(WRR, 2013, A stochastic model for estimating groundwater and contaminant discharges from fractured rock passive flux meter measurements)

# Observed Data

- Number of intersections **N**
- For each intersection:
  - Trace location and orientation
  - Flux per unit trace length
- Trace density  $\lambda$  and trace length distribution **pdf( $\tau$ )** remain **UNKNOWN**
- **HOWEVER**, both **N** and discharge **Q** are proportional to the product  $\lambda\mu_\tau$  ( $\mu_\tau$  ... mean trace length)
- **Trace density and length properties do not have to be explicitly known for discharge estimation !!!**



# Discharge Estimator $Q^*$

$$Q^* = A_T \sum_{k=1}^{N_{well}} \frac{\omega_k}{L_k} \sum_{j=1}^{N_k} \frac{q_j^*}{\cos \theta_j}$$

$A_T$  ... Area of control plane

$N_{well}$  ... Number of borings in control plane

$N_k$  ... Number of intersections on the k-th boring

$L_k$  ... Lengths of k-th boring

$\omega_k$  ... Weight of k-th boring (for irregular separation distances)

$\theta_j$  ... Orientation angle of j-th intersection with k-th boring

$q_j^*$  ... Measured flux at j-th intersection with k-th boring

# Upper Uncertainty Bound

$$CV_e^2 \leq \frac{1 + CV_{\rho^*}^2 + \frac{CV_{\rho^*}^4}{4}}{N_{total}}$$

**CV<sub>e</sub>** ... Error coefficient of variation

**N<sub>total</sub>** ... Total number of intersections over all borings

**CV<sub>ρ\*</sub>** ... CV of measured fluxes after simple transformation

- Ex.: For **CV<sub>ρ\*</sub> ≈ 1** and **N<sub>total</sub> ≈ 50** we get **CV<sub>e</sub> ≤ 0.21**
- Device for measuring fracture fluxes is currently being field tested → presentation **TOMORROW AFTERNOON**

# Summary

- Contaminant mass discharges are useful parameters in site characterization and decision making, **IF** uncertainty is quantified
- Local flux measurements heavily simplify and improve discharge and uncertainty estimation (no flow and transport modeling required)
- Approaches presented for porous and fractured rock aquifers
- Both require sufficient data to reliably represent flux heterogeneity across control plane

**Thank you!**

**QUESTIONS ?**